

Subdiffusive flow in a composite medium with a communicating (absorbing) interface

Rajagopal Raghavan^{1,*} and Chih-Cheng Chen²

¹The Raghavan Group, PO Box 52756, Tulsa, OK 74152, USA

²Kappa Engineering, 11767 Katy Freeway, #500, Houston, TX 77079, USA

Received: 23 October 2019 / Accepted: 2 March 2020

Abstract. Two-dimensional subdiffusion in media separated by a partially communicating interface is considered. Starting with the appropriate Green’s functions, solutions are developed in terms of the Laplace transformation reflecting two circumstances at the interface: situations where there is perfect contact and situations where the interface offers a resistance. Asymptotic solutions are derived; limiting forms of the expressions reduce to known solutions for both classical diffusion and subdiffusion. Specifics are analyzed in depth with reference to flow in porous media with potential applications to the evaluation of the role of subsurface faults and flow in fractured rocks. Characteristics of the derivative responses are documented extensively as they are the linchpin for evaluation of pressure tests. Results given here may be used for evaluation at the Theis (1935; *Eos Trans. AGU* **2**, 519–524) scale along with geological and geophysical properties, and production statistics. Yet a subdiffusive model does not imply a single value for properties. The method presented here may be extended to multiple contiguous media and to subdiffusive transport in many contexts (complex wellbores such as inclined, fractured and horizontal wells, situations such as sequestration, production of geothermal systems, etc.).

Nomenclature

A_j	Constants; see equations (28)–(30)	p	Pressure [M/L/T ²]
a	Defined in equation (7)	p_i	Initial pressure [M/L/T ²]
B	Formation volume factor [L^3/L^3]	p'	Logarithmic derivative [M/L/T ²]
c_t	Total compressibility [L T ²]	q	Rate [L^3/T]
d	Dimension	\Re	Real part of a number
$E_{\alpha,\beta}(z)$	Mittag-Leffler function; see equation (9)	r	Distance [L]
$f_j(\varphi)$	See equations (18) and (19)	s	Laplace transform variable [1/T]
h	Thickness [L]	t	Time [T]
$H_{p,q}^{m,n} \left[x \begin{matrix} (a_1, A_1), \dots, (a_p, A_p) \\ (b_1, B_1), \dots, (b_q, B_q) \end{matrix} \right]$	Fox function	T_D	See equation (46)
		T_r	See equation (31)
		$\mathbf{v}(\mathbf{x}, t)$	Flux [L/T]
		$W(\alpha, \beta; t)$	Wright function
		W	Width [L]
$K_\nu(z)$	Modified Bessel function; second kind of order ν	\mathbf{x}	Coordinates [L]
k	Permeability [L ²]	z	A complex function
\tilde{k}	Subdiffusive permeability; see equation (11) [L ² /T ^{1-α}]	α	Exponent; see equation (10)
ℓ	Reference length [L]	$\gamma_j(\mathbf{x}; t)$	Source functions; equations (18) and (19)
L	Distance between the interface and the well [L]	$\Gamma(\cdot)$	Gamma function
		γ	Euler’s constant (0.5772...)
		η	Diffusivity; various
		$\tilde{\eta}$	“Diffusivity”; see equation (20) [L ² /T ^{α}]
		η_D	See equation (48)
		λ	Mobility [L T/M]

* Corresponding author: raghavan.raj@gmail.com

$\tilde{\lambda}$	Equation (11)
μ	Viscosity [M/L/T]
ϕ	Porosity [L ³ /L ³]
$\psi(x)$	Digamma function
Υ	Contact resistance; see equation (16)
Υ_D	See equation (47)
Υ_r	See equation (32)

Subscripts

D	Dimensionless
F	Fault
i	Initial
j	Zone with distinct properties
w	Well, wellbore

Superscripts

–	Laplace transform
---	-------------------

1 Introduction

Interfaces are ubiquitous in many fields of endeavor; extensive literature exists in a range of technologies (biophysics, geophysics, energy recovery, tomography, acoustics and electromagnetism to mention a few). Developments in heat transmission made through Green’s function techniques are often the starting point in many studies; see Carslaw and Jaeger (1959), in both one and two dimensions for single and multilayered media. Often point- or line-source solutions are employed as in Mandelis *et al.* (2001) and Shendeleva (2004). More recently, subdiffusive flow in linear, composite media have been explored (Kosztowicz, 2017; Povstenko, 2013; Povstenko and Kyrylych, 2019; Raghavan and Chen, 2018). The motivation for such studies, as in all other cases, is to evaluate the consequences of finite speeds of propagation of the transient in linear, composite systems of the form

$$\langle r^2 \rangle \sim t^a, \tag{1}$$

with r being the distance, t being time and where $a \neq 1$ is a constant. Equation (1) for $a < 1$, namely subdiffusion, has drawn much interest in recent times because of diffusion in fractal structures (Gefen *et al.*, 1983; Metzler *et al.*, 1994; Nigmatullin, 1984a, b), although interest in finite speeds of propagation has been discussed for a long time; see Gurtin and Pipkin (1968). More recent examples of subdiffusion from an application vantage point particularly from the perspective of flow in porous media, the focus of this work, may be found in Caputo (1999), Jourde *et al.* (2002), Cortis and Knudby (2006) and Raghavan (2011). In this work we discuss influence of an interface under subdiffusion in 2D as no such studies exist. We examine the problem in the context of application to earth sciences such as the flow of groundwater and extraction of hydrocarbons by considering the interface in the form of a partial

hydrologic barrier. Partial hydrologic barriers are of particular interest as they could serve as conduits at times or seals in other times (Bense and Person, 2006; Miller, 2005; Raghavan and Ozkan, 2011). There is evidence in the earth science literature that the permeability changes significantly around faults and that damaged regions with cracks, fissures and cervices abound around most faults. As a result, scaling laws that systematically reflect the damage around faults have been developed (Caine *et al.*, 1996; Evans, 1988; Mitchell and Faulkner, 2009; Savage and Brodsky, 2011; Scholz *et al.*, 1993). Liesegang bands, regions of reduced permeability around fracture-matrix interfaces, are modeled in classical situations as skin regions (Prats and Raghavan, 2013, 2014) (ignoring scaling laws), and the phenomenon is known as interporosity skin (Cinco-Ley *et al.*, 1985); photographs of Liesegang bands may be found in Fu *et al.* (1994), Sharp *et al.* (1996) and Neville *et al.* (1998). As we employ the method of sources and sinks, our results may be directly transposed to apply to other disciplines by a change in nomenclature. The source functions to determine the needed solutions are constructed in the manner of Raghavan and Ozkan (1994) in terms of the Laplace transformation with the method proposed in Sommerfeld (1909). The effectiveness of his method is well-attested by the number of researchers who have employed Sommerfeld’s technique in a wide range of disciplines such as electromagnetic radiation (Lindell and Alanen, 1984; Lindell *et al.*, 1986), geophysics (Hänninen *et al.*, 2002) and flow in porous media (Raghavan, 2010) in addition to heat conduction (Shendeleva, 2004).

We presume two options to represent conditions at the interface: perfect contact or a steady-state resistance as modeled in Carslaw and Jaeger (1959). The latter observation implies that the resistance either causes or produces no transients. We may readily expand the solutions derived for other well configurations if desired; see Raghavan (2012) or address anisotropy as we arrive at expressions for the instantaneous source function, $\gamma(\mathbf{x}; t)$, very similar in structure to that used to represent various forms of linear barriers with image wells of the form described in Tualle *et al.* (2000):

$$\gamma(x, y; t) = \Pi \otimes S + \Pi \otimes S', \tag{2}$$

where the symbol S represents the source, the symbol S' represents the image sources that are distributed to reflect the interfaces, the symbol Π is the solution for a source placed in an infinite medium with the properties of the region where the source S is located and the symbol \otimes represents convolution in both space and time.

1.1 The Mathematical model

Transient diffusion in a porous rock of thickness, h , structured in the form of a composite system with an interface located along the plane $y = 0$ and extending to infinity in both the x and y directions is considered. The influence of the distinct properties of the rock in a connected domain, Ω , in \mathcal{R}^2 on either side of the interface is examined. A contact resistance at the interface governs fluid movement and

the system is produced through a line-source well located at the point $(0, y')$. In and of itself the interface has no storage capacity; consequently no transient effects are produced as the result of its existence. As is usual we consider a slightly compressible liquid of compressibility, c , and viscosity μ . Initially, throughout the system the pressure, $p(x, t)$, is assumed to be uniform at the value of p_i . The transient diffusion equation for this configuration is

$$\nabla \cdot [v(\mathbf{x}, t)] = \phi(\mathbf{x})c_t \frac{\partial p(\mathbf{x}, t)}{\partial t}, \quad (3)$$

where the symbols $v(\mathbf{x}, t)$ and $\phi(\mathbf{x})$ represent the flux and porosity, respectively. This expression is solved for the following constraints: Because the system is quiescent initially and also infinite in areal extent we require,

$$p(\mathbf{x}; 0) = p_i, \quad \mathbf{x} \in \Omega; \quad (4)$$

and also

$$p(\{x, y\} \rightarrow \infty; t) = p_i. \quad (5)$$

Assuming flow to a line-source well, the condition at a wellbore for a specified withdrawal rate, q , is

$$\lim_{r \rightarrow 0} [2\pi h r v(r)] = qB, \quad (6)$$

where B is the formation volume factor, and r is the distance between the point (x, y) and the well location $(0, y')$.

To satisfy the conditions at the interface, $y = 0$, we require that the pressure drop across the interface be proportional to the normal flux and also the normal fluxes be continuous both for all time, t (Yaxley, 1987); thus, we require that

$$v_{x_1}(x, y; t) = a[p_1(x, y; t) - p_2(x, y; t)], \quad y = 0, \quad (7)$$

and

$$v_{x_1}(x, y; t) = v_{x_2}(x, y; t), \quad y = 0, \quad (8)$$

where the symbol $a = k_F/(\mu w_F)$ is a constant and the subscript ‘‘F’’ signifies the fault. If a were 0, then the condition specified in equation (7) reflects perfect contact and thus the continuity in pressure for all time, t . In essence, steady-state Darcy flow is assumed across the fault and the fault causes no transients in of itself as it contains no pore volume. The subscripts 1 and 2 in the above expressions represent the regions on either side of the interface; henceforth, we presume Zone 1 represents the region where the well is located. We have treated the existence of the discontinuity at $y = 0$ in a manner very similar to that of a resistance discussed in Carslaw and Jaeger (1959).

In the following we work in terms of the Laplace transformation and use the method of sources and sinks; see Raghavan and Ozkan (1994). As already mentioned, the influence of the interface is incorporated by image sources located symmetrically about the interface with the resulting pressure distribution determined through convolution.

1.2 The Darcy law, subdiffusive flow

The experiments that Darcy (1856) conducted to arrive at the flux law that the superficial velocity is proportional to the potential difference by considering the flow of water in porous columns may be derived from general energy considerations as shown in Miller (1954). For his experiments, a number of considerations are implicit as set out in Childs (1969). Specifically, as noted in Raghavan (1993) five factors are inherent: (i) the porous medium is an uniform body as it is appreciably larger than the pores that constitute the porous rock, (ii) the structure of the porous rock and the potentials considered are notional surfaces; in essence, the different cross-sectional elements that permit fluid flow are so alike that the fluxes are identical and the potential surfaces reflect the overall flux, (iii) the flux is inversely proportional to the length of the porous column, (iv) motion is independent of time (Philip, 1957), and (v) inertia terms are nonzero but are of little consequence (Lindquist, 1933). In the context of random walks, the flux law such as the Darcy law is equivalent to a particle traveling a fixed distance, Δx , over a fixed time, Δt .

Should the structure of the porous medium be considered to be a fractal porous object (Gefen *et al.*, 1983; O’Shaughnessy and Procaccia, 1985a, b), then the nature of the flux law changes; both permeability and porosity are power-law functions of the Euclidean dimension, d , the fractal or Hausdorff dimension d_f , and the random walk dimension or anomalous diffusion coefficient d_w (Chang and Yortsos, 1990) because the diffusivity of the porous rock, η , follows the relation $\eta \sim r^{-(d_w-2)}$, where r is distance; see Gefen *et al.* (1983). Other fractal models such as those proposed in Metzler *et al.* (1994), though never specifically noted, require a different flux law. The easiest way to model their work is to consider a time dependent diffusivity which leads to the fractional differential equation proposed by them.

Anomalous diffusion, the focus of this work, is best described in terms of a general random walk, namely a Continuous-Time Random Walk (CTRW) proposed by Montroll and Weiss (1965); also see Noetinger and Estebenet (2000) and Henry *et al.* (2010). Unlike standard random walks that use a fixed step length, Δx , at intervals that are separated through a fixed time interval, Δt , the CTRW formulation uses an i.i.d. (independent, identically distributed) random variable for increments of the step length, Δx , and a waiting (pausing) time function, $\psi(t)$, that determines the time between jumps. Kenkre *et al.* (1973) show that a CTRW formulation leads to an equation (part of the family of master equations) that is equivalent to the Montroll–Weiss formulation. Hilfer and Anton (1995) connect the CTRW formulation to fractional differential equations through the waiting time function $\psi(t) = (t^\omega/C)E_{\omega,\omega}(-t^\omega/C)$, where $E_{\omega,\omega}(x)$ is the Mittag-Leffler function discussed in many texts (Erdelyi *et al.*, 1955; Mainardi, 2010; Povstenko, 2015; Uchaikin, 2013):

$$E_{\alpha,\beta}(z) = \sum_{n=0}^{\infty} \frac{z^n}{\Gamma(\beta + \alpha n)}, \quad \alpha, \beta > 0, \quad z \in C, \quad (9)$$

where C represents the complex plane and where $\Gamma(\cdot)$ is the Gamma function. The symbol ω is a constant that depends on the properties of the fractal object; with $0 \leq \omega \leq 1$. In a similar manner, Henry *et al.* (2010) demonstrate that if we were to use two probability density (pdfs) functions, one a step length density function, Δx , that is an even function and the other a power-law waiting time function, $\psi(t)$, that is scale invariant with temporal memory, then such an option also leads to a fractional flux law. For instance, the Pareto waiting time density (see Feller, 1971) with a mean waiting time that is infinite could be used for, $\psi(t)$.

The CTRW formulation leads to a flux law of the following form:

$$v(\mathbf{x}, t) = -\tilde{\lambda} \frac{\partial^{1-\alpha}}{\partial t^{1-\alpha}} [\nabla p(\mathbf{x}, t)], \quad (10)$$

with

$$\tilde{\lambda} = \frac{\tilde{k}}{\mu}. \quad (11)$$

Here α is the subdiffusion coefficient with $0 < \alpha < 1$, μ is the fluid viscosity, \tilde{k} is the permeability corresponding to α and $\partial^\alpha f(t)/\partial t^\alpha$, is the fractional derivative defined in Caputo (1967) sense; thus

$$\frac{\partial^\alpha}{\partial t^\alpha} f(t) = \frac{1}{\Gamma(1-\alpha)} \int_0^t dt (t-t')^{-\alpha} \frac{\partial}{\partial t'} f(t'). \quad (12)$$

For $\alpha = 1$, equation (10) reverts to the classical Darcy equation. The Laplace transformation of the Caputo fractional operator, ${}_0^c D_t^\alpha f(t)$, is given by

$$\int_0^\infty e^{-st} {}_0^c D_t^\alpha f(t) dt = s^\alpha f(s) - s^{\alpha-1} f(0), \quad 0 \leq \alpha < 1. \quad (13)$$

Implicit in this definition is the fact that the flux at any instant, t , may not be obtained directly from the instantaneous pressure distribution, $p(\mathbf{x}, t)$.

For subdiffusive flows, fractional flux laws of the form proposed here have been used for many years in a variety of contexts concerning diffusion: in biophysics (Magin *et al.*, 2013), fractal structures (Dassas and Duby, 1995; Metzler *et al.*, 1994), flow in porous media (Albinali *et al.*, 2016; Lē Mehautē and Crepy, 1983; Nigmatullin, 1984a, b; Płociniczak, 2015; Raghavan, 2011; Raghavan and Chen, 2017; Su, 2014; Su *et al.*, 2015), contaminant transport (Fomin *et al.*, 2011; Kim *et al.*, 2015; Molz *et al.*, 2002; Suzuki *et al.*, 2016), heat transmission (Angulo *et al.*, 2000; Gurtin and Pipkin, 1968; Moodie and Tait, 1983; Norwood, 1972). Experimental aspects are discussed in Tao *et al.* (2016), Zhokh and Strizhak (2018, 2019) and Yanga *et al.* (2019). Additional options are available to modify equation (10) should one address phenomena such as truncation (Odling *et al.*, 1999) to model structures that are discontinuous resulting in the variation of permeability around faults. Also scalings that result in power law variations yield cutoffs which may result in changes in flow

regimes (classical to subdiffusion). In this work we, however, limit our discussion to equation (10).

While on the subject of pdfs, we simply note that a standard random walk may be modeled by considering the Markovian exponential waiting time density (memorylessness; see Feller, 1971).

1.3 Formulation

Using $\Delta p_j(\mathbf{x}, t) = p_i - p_j(\mathbf{x}, t)$, where the subscript $j = 1, 2$ reflects Zone 1 and Zone 2, respectively, we may combine equations (3) and (10) to obtain the transient diffusion partial differential equations for subdiffusive flow in Zones 1 and 2, respectively, to be

$$\nabla \cdot [\tilde{\lambda}_j \nabla \Delta p_j(\mathbf{x}, t)] = \phi_j c_{ij} \frac{\partial^{z_j}}{\partial t^{z_j}} [\Delta p_j(\mathbf{x}, t)]. \quad (14)$$

Similarly, equations (6)–(8) lead, respectively, to

$$\left[r \frac{\partial^{1-\alpha_1}}{\partial t^{1-\alpha_1}} \left(\frac{\partial p}{\partial r} \right) \right]_{r=0} = \frac{qB\mu}{2\pi \tilde{k}_j h}, \quad t > 0, \quad (15)$$

$$-\Upsilon \frac{\partial^{1-\alpha_1}}{\partial t^{1-\alpha_1}} \left[\frac{\partial \Delta p_1}{\partial x} (x, y; t) \right] = \Delta p_2(x, y; t) - \Delta p_1(x, y; t), \quad y = 0, \quad (16)$$

where $\Upsilon = \tilde{k}_1 w_F / k_F$, and

$$\tilde{\lambda}_1 \frac{\partial^{1-\alpha_1}}{\partial t^{1-\alpha_1}} \left(\frac{\partial \Delta p_1}{\partial x} \right) = \tilde{\lambda}_2 \frac{\partial^{1-\alpha_2}}{\partial t^{1-\alpha_2}} \left(\frac{\partial \Delta p_2}{\partial x} \right), \quad y = 0. \quad (17)$$

1.4 General solution

Instantaneous line sources of the kind needed in this study may be constructed by choosing expressions to satisfy the basic partial differential equation and the auxiliary conditions; for line sources considered in this work one such expression is the modified Bessel function of the second kind of order 0, $K_0(x)$, that reflects the free subdiffusive propagator, $\mathcal{G}(\mathbf{r}; t)$, in \mathcal{R}^2 . Accordingly, for Zone 1 and Zone 2 we write instantaneous solutions, $\gamma_j(x, y)$; $j = 1, 2$, respectively, in terms of their Laplace transformation, $\bar{\gamma}_j(x, y)$ with the source located at $(0, y')$ to be

$$\bar{\gamma}_1(x, y) = \frac{1}{2\pi s^{1-\alpha_1} \tilde{\eta}_1} K_0 \left(\sqrt{s^{\alpha_1} / \tilde{\eta}_1} r_1 \right) + \int_0^\infty f_1(\varphi) e^{-y\varphi_1} \cos(\varphi x) d\varphi, \quad (18)$$

and

$$\bar{\gamma}_2(x, y) = \int_0^\infty f_2(\varphi) e^{y\varphi_2} \cos(\varphi x) d\varphi, \quad (19)$$

where $\tilde{\eta}_j$, the “diffusivity” of the medium, is given by

$$\tilde{\eta}_j = \frac{\tilde{k}_j}{\phi_j c \mu}, \quad (20)$$

and

$$r_1^2 = x^2 + (y - y')^2 \tag{21}$$

with the symbol q_j given by

$$q_j = \sqrt{\phi^2 + \frac{s^{z_j}}{\tilde{\eta}_j}} \tag{22}$$

The motivation for the above expressions is the observation given in Erdélyi *et al.* (1954, see p. 17); Gradshteyn and Ryzhik (1994, p. 532; 3.961–3.9612) that the integral representation of $K_0(x)$ is

$$K_0 \left[a(x^2 + y^2)^{\frac{1}{2}} \right] = \int_0^\infty \exp \left[-y(u^2 + a^2)^{\frac{1}{2}} \right] \frac{\cos(xu)}{(u^2 + a^2)^{\frac{1}{2}}} du \tag{23}$$

In addition, the first term is the Green’s function for a medium that is infinite in lateral extent. Our procedure, first proposed by Sommerfeld (1909) has been used by many; see, for example, Mandelis *et al.* (2001), Shendeleva (2004) and Raghavan (2010).

On using equations (16) and (17) to solve for $f_1(\phi)$ and $f_2(\phi)$, we obtain the following expressions:

$$\begin{aligned} \bar{\gamma}_1(\mathbf{x}) = & \frac{1}{2\pi s^{1-z_1} \tilde{\eta}_1} \left[K_0 \left(\sqrt{s^{z_1} / \tilde{\eta}_1} r_1 \right) \right. \\ & \left. + \int_0^\infty A_1 e^{-q_1(y+y')} \cos(\phi x) d\phi \right], \end{aligned} \tag{24}$$

or

$$\begin{aligned} \bar{\gamma}_1(\mathbf{x}) = & \frac{1}{2\pi s^{1-z_1} \tilde{\eta}_1} \left[K_0 \left(\sqrt{s^{z_1} / \tilde{\eta}_1} r_1 \right) - K_0 \left(\sqrt{s^{z_1} / \tilde{\eta}_1} r_2 \right) \right. \\ & \left. + 2 \int_0^\infty A_2 e^{-q_1(y+y')} \cos(\phi x) d\phi \right], \end{aligned} \tag{25}$$

and,

$$\bar{\gamma}_2(\mathbf{x}) = \frac{1}{\pi s^{1-z_1} \tilde{\eta}_1} \int_0^\infty A_3 e^{(q_2 y - q_1 y')} \cos(\phi x) d\phi, \tag{26}$$

where

$$r_2^2 = x^2 + (y + y')^2. \tag{27}$$

The functions A_1 , A_2 and A_3 are given, respectively, by

$$A_1 = \frac{q_1 - (1 - \Upsilon_r q_1) T_r q_2}{q_1 [q_1 + (1 + \Upsilon_r q_1) T_r q_2]}, \tag{28}$$

$$A_2 = \frac{1 + \Upsilon_r T_r q_2}{q_1 + (1 + \Upsilon_r q_1) T_r q_2}, \tag{29}$$

and

$$A_3 = \frac{1}{q_1 + (1 + \Upsilon_r q_1) T_r q_2}. \tag{30}$$

Here

$$T_r = \frac{T_2}{T_1} s^{z_1 - z_2}, \tag{31}$$

$$\Upsilon_r = \frac{\tilde{k}_1 w_F}{k_F} s^{1-z_1} \tag{32}$$

and $T_j = \tilde{k}_j h$ with the source functions on hand we may write the pressure distributions through the procedure followed in Raghavan and Ozkan (1994). The pressure in terms of the Laplace transform of $\Delta p(x, y, z, t)$ (Raghavan and Ozkan, 1994) is

$$\overline{\Delta p}(\mathbf{x}) = \frac{1}{\phi c} \int_S \bar{q}(x', y') \bar{\gamma}(x - y', y - y') dS'. \tag{33}$$

Here (x', y') is the location of the source, (x, y) any location in the porous rock and the strength of the continuous line-source is $\bar{q}(x', y'; t)/(\phi c)$ with dS' being the element (a line or surface) through which fluid is extracted. For an instantaneous line-source with the source strength being constant, we have

$$\overline{\Delta p}(\mathbf{x}) = \frac{1}{s} \frac{\tilde{q}}{\phi c} \int_S \bar{\gamma}(x - x', y - y', z - z') dS'. \tag{34}$$

Thus on using equation (34) and the expressions for $\bar{\gamma}_j(x, y)$ derived above, the expressions for $\Delta p_j(x, y; t)$ in terms of their respective Laplace transformations, $\overline{\Delta p}_j(x, y)$, are

$$\begin{aligned} \overline{\Delta p}_1(x, y) = & \frac{1}{s^{2-z_1}} \frac{q_w B \mu}{2\pi T_1} \left[K_0 \left(\sqrt{s^{z_1} / \tilde{\eta}_1} r_1 \right) \right. \\ & \left. + \int_0^\infty A_1 e^{-q_1(y+y')} \cos(\phi x) d\phi \right], \end{aligned} \tag{35}$$

or

$$\begin{aligned} \overline{\Delta p}_1(x, y) = & \frac{1}{s^{2-z_1}} \frac{q_w B \mu}{2\pi T_1} \left[K_0 \left(\sqrt{s^{z_1} / \tilde{\eta}_1} r_1 \right) - K_0 \left(\sqrt{s^{z_1} / \tilde{\eta}_1} r_2 \right) \right. \\ & \left. + 2 \int_0^\infty A_2 e^{-q_1(y+y')} \cos(\phi x) d\phi \right], \end{aligned} \tag{36}$$

in Zone 1 ($y > 0$), and

$$\overline{\Delta p}_2(x, y) = \frac{1}{s^{2-z_1}} \frac{q_w B \mu}{\pi T_1} \int_0^\infty A_3 e^{(q_2 y - q_1 y')} \cos(\phi x) d\phi, \tag{37}$$

in Zone 2 ($y < 0$).

1.5 Solutions for perfect contact

To proceed further, we define dimensionless variables for pressure, $p_D(\mathbf{x}_D; t_D)$, time, t_D , and distance, \mathbf{x}_D , respectively, as follows:

$$p_D(\mathbf{x}_D; t_D) = \frac{2\pi \tilde{k}_1 h}{q_w B \mu} \left(\frac{\tilde{\eta}_1}{\ell^2} \right)^{\frac{1-z_1}{z_1}} \Delta p(\mathbf{x}; t), \tag{38}$$

$$t_D = \frac{\tilde{\eta}_1}{\ell^2} t^{\alpha_1}, \tag{39}$$

and

$$\mathbf{x}_D = \frac{\mathbf{x}}{\ell}, \tag{40}$$

where ℓ is the reference length.

Using these definitions, we obtain the following expression for the pressure distributions, $\bar{p}_{Dj}(\mathbf{x}_D)$; $j = 1$ and 2 , in terms of the Laplace transformation for Zones 1 and 2 to be, respectively,

$$\begin{aligned} \bar{p}_{D1}(\mathbf{x}_D) = & \left(\frac{\tilde{\eta}_1}{\ell^2}\right)^{\frac{1-\alpha_1}{\alpha_1}} \frac{1}{s^{2-\alpha_1}} \left[K_0\left(\sqrt{s^{\alpha_1} \ell / \tilde{\eta}_1} r_{D1}\right) - K_0\left(\sqrt{s^{\alpha_1} \ell^2 / \tilde{\eta}_1} r_{D2}\right) \right. \\ & \left. + 2 \int_0^\infty A_2 e^{-q_1(y_D+y'_D)} \cos(\varphi x_D) d\varphi \right], \end{aligned} \tag{41}$$

for $y_D > 0$ and

$$\bar{p}_{D2}(\mathbf{x}_D) = \left(\frac{\tilde{\eta}_1}{\ell^2}\right)^{\frac{1-\alpha_1}{\alpha_1}} \frac{2}{s^{2-\alpha_1}} \int_0^\infty A_3 e^{(q_2 y_D - q_1 y'_D)} \cos(\varphi x_D) d\varphi, \tag{42}$$

for $y_D < 0$. If we consider the integrals in equations (36) and (37), we see that the response at any point M for the interface located at a distance, L from the well is a function of Υ , $\tilde{\eta}_r$ and T_r where

$$\eta_r = \frac{\tilde{\eta}_1}{\tilde{\eta}_2}. \tag{43}$$

But the A_j 's; that is, both q_j 's, and the three variables, T_r , $\tilde{\eta}_r$, and Y_r , must, of course, be rescaled. The scalings are:

$$T_r = \frac{T_2}{T_1} s^{\alpha_1 - \alpha_2} = T_D \left(\frac{\ell^2 s^{\alpha_1}}{\tilde{\eta}_1}\right)^{1-\alpha_D}, \tag{44}$$

$$\Upsilon_r = \Upsilon_D \left(\frac{\ell^2 s^{\alpha_1}}{\tilde{\eta}_1}\right)^{\frac{1-\alpha_1}{\alpha_1}}, \tag{45}$$

with

$$T_D = \frac{T_2}{T_1} \left(\frac{\tilde{\eta}_1}{\ell^2}\right)^{(1-\alpha_D)}, \tag{46}$$

$$\Upsilon_D = \frac{\tilde{k}_1 w_F}{\tilde{k}_F \ell} \left(\frac{\tilde{\eta}_1}{\ell^2}\right)^{\frac{1-\alpha_1}{\alpha_1}}, \tag{47}$$

$$\eta_D = \frac{\tilde{\eta}_1^{2D}}{\tilde{\eta}_2} \ell^{2(1-\alpha_D)}, \tag{48}$$

and

$$\alpha_D = \frac{\alpha_2}{\alpha_1}. \tag{49}$$

Scalings involving the ratio of α_j 's enter into the picture because we consider subdiffusive flow; in accordance with equation (10); the dimension of \tilde{k} is a function of α . These scalings also follow if the partial differential equations are considered. The groupings T_D and η_D may be replaced by T_{D1} and $(\phi c_i h)_D$ if desired, where

$$(\phi c_i h)_D = \frac{(\phi c_i h)_1}{(\phi c_i h)_2}. \tag{50}$$

A number of approximate solutions may be derived if $q_1 = q_2$, which implies $\tilde{\eta}_1$ and $\tilde{\eta}_2$ and also $\alpha_1 = \alpha_2$. In the following we use $\ell = r_w$, where r_w is the well radius. In this situation for Zone 1 and Zone 2, respectively, we have

$$\begin{aligned} \bar{p}_{D1}(\mathbf{x}_D) = & \left(\frac{\tilde{\eta}_1}{r_w^2}\right)^{\frac{1-\alpha_1}{\alpha_1}} \frac{1}{s^{2-\alpha_1}} \left[K_0\left(\sqrt{s^{\alpha_1} r_w^2 / \tilde{\eta}_1} r_{D1}\right) \right. \\ & \left. - \frac{1 - T_D}{1 + T_D} K_0\left(\sqrt{s^{\alpha_1} r_w^2 / \tilde{\eta}_1} r_{D2}\right) \right], \end{aligned} \tag{51}$$

and

$$\bar{p}_{D2}(\mathbf{x}_D) = \left(\frac{\tilde{\eta}_1}{r_w^2}\right)^{\frac{1-\alpha_1}{\alpha_1}} \frac{2}{(1 + T_r)} \frac{1}{s^{2-\alpha_1}} K_0\left(\sqrt{s^{\alpha_1} r_w^2 / \tilde{\eta}_1} r_{D1}\right). \tag{52}$$

The right-hand sides of equations (51) and (52) may be inverted with the expression

$$2L^{-1}[s^{-\varrho} K_\nu(zs^\zeta)] = t^{\varrho-1} H_{1,2}^{2,0} \left[\frac{z^2 t^{-2\zeta}}{4} \left| \begin{matrix} \varrho, 2\zeta \\ (\frac{\varrho}{2}, 1), (-\frac{\varrho}{2}, 1) \end{matrix} \right. \right], \tag{53}$$

where the symbol $H_{p,q}^{m,n} \left[x \left| \begin{matrix} (a_1, A_1), \dots, (a_p, A_p) \\ (b_1, B_1), \dots, (b_q, B_q) \end{matrix} \right. \right]$ is the Fox function or the H-function, $\Re(\varrho) > 0$, $\Re(z^2) > 0$, $\Re(s) > 0$ as noted in Saxena *et al.* (2006), and thus we have the following result for $p_{D1}(\mathbf{x}_D; t_D)$, namely

$$\begin{aligned} p_{D1}(\mathbf{x}_D; t_D) = & \frac{1}{2} t_D^{\frac{1-\alpha_1}{\alpha_1}} \\ & \times \sum_{j=0}^{j=1} \left\{ \left(\frac{1 - T_D}{1 + T_D}\right)^j H_{1,2}^{2,0} \left[\frac{r_{Dj+1}^2}{4t_D} \left| \begin{matrix} 2-\alpha_1, \alpha_1 \\ (0,1), (0,1) \end{matrix} \right. \right] \right\}, \end{aligned} \tag{54}$$

and for $p_{D2}(\mathbf{x}_D; t_D)$, namely

$$p_{D2}(\mathbf{x}_D; t_D) = \frac{1}{(1 + T_r)} t_D^{\frac{1-\alpha_1}{\alpha_1}} \left\{ H_{1,2}^{2,0} \left[\frac{r_{D1}^2}{4t_D} \left| \begin{matrix} 2-\alpha_1, \alpha_1 \\ (0,1), (0,1) \end{matrix} \right. \right] \right\}. \tag{55}$$

Although many discussions of Fox functions are available; see Mathai and Saxena (1978) and Mainardi *et al.* (2005), such functions are best evaluated through packages such as Mathematica. Also, if $j = 0$; that is, one well in a system that is infinite in its extent, equation (55) reduces to the expression, as expected, in Raghavan and Chen (2019).

Along the same lines, we may obtain a long-time approximation for this specific situation by considering

the following expression for $K_0(x)$ as $x \rightarrow 0$ (implicitly we consider situations where $s \rightarrow 0$), namely using

$$K_0(z) = -\left(\ln \frac{z}{2} + \gamma\right) + \frac{z^2}{4} \left[1 - \left(\gamma + \ln \frac{z}{2}\right)\right] + O(z^4 \ln z), \quad (56)$$

where γ is Euler's constant. The long-time pressure response, $p_D(\mathbf{x}_D; t_D \rightarrow \infty)$, may be obtained by using the first two terms of the above expression, as well as the result of Oberhettinger and Badii (1973)

$$s^{-\nu} \ln s = [\Gamma(\nu)]^{-1} t^{\nu-1} [\psi(\nu) - \ln t], \quad (57)$$

where $\psi(\cdot)$ is the Digamma function, and we therefore have

$$p_{D_1}(\mathbf{x}_D, t_D) = \frac{1}{2} \frac{1}{\Gamma(2-\alpha_1)} t_D^{\frac{1-\alpha_1}{\alpha_1}} \left\{ \left[\ln \frac{4t_D}{e^{2\gamma} r_{D_1}^2} - \alpha_1 \psi(2 - \alpha_1) \right] + \left(\frac{1-T_D}{1+T_D} \right) \left[\ln \frac{4t_D}{e^{2\gamma} r_{D_2}^2} - \alpha_1 \psi(2 - \alpha_1) \right] \right\}. \quad (58)$$

Also, the logarithmic derivative, $p'_{D_1}(\mathbf{x}_D, t_D)$, is given by the expression

$$p'_{D_1}(\mathbf{x}_D, t_D) = \frac{1}{2} \frac{1}{\Gamma(1-\alpha_1)} t_D^{\frac{1-\alpha_1}{\alpha_1}} \left\{ \left[\ln \frac{4t_D}{e^{2\gamma} r_{D_1}^2} - \alpha_1 \psi(1 - \alpha_1) \right] + \left(\frac{1-T_D}{1+T_D} \right) \left[\ln \frac{4t_D}{e^{2\gamma} r_{D_2}^2} - \alpha_1 \psi(1 - \alpha_1) \right] \right\}. \quad (59)$$

Focussing on the first term of equation (58) for now and assuming $T_D = 1$, we note that for evaluating fractured reservoirs, interestingly, Bernard *et al.* (2006) propose that responses be evaluated as a function of $t \ln t$; this work perhaps gives a foundation for their empirical recommendation. Both expressions, equations (55) and (58) for $\alpha = 1$ in an infinitely large system ($T_D = 1$), respectively, correspond to the Theis (1935) solution and the semi-logarithmic approximation of the Theis (1935) solution, namely the Cooper and Jacob (1946) approximation. Along the same lines if both α_j 's are unity, then both these expressions correspond to the ones given in Raghavan (2010). Most importantly, it is to be noted that unlike classical diffusion, for subdiffusive flows the influence of Zone 1 never stabilizes; that is, one must consider transient effects caused by Zone 1 for all times. This difference is exceedingly important and is a direct result of the observation made earlier that the flux, $v(\mathbf{x}, t)$, is not obtained from the instantaneous pressure distribution, $p(\mathbf{x}, t)$.

At long enough times the slope of the second line would reflect $2/(1 + T_D)$, if both lines corresponding to Zones 1 and 2 were evident, which would require values of L to be neither too small nor too large, where L is the distance between the interface and well. Then $(1 - T_D)/(1 + T_D)$ may be estimated by extending the line corresponding to Zone 1 and determining the separation between the two lines.

Unfortunately, it does not appear to be possible to extend the above results for $\tilde{\eta}_2$ different from $\tilde{\eta}_1$ at this time. Although this is the case, we know that for classical diffusion ($\alpha_j = 1$) the influence of $\tilde{\eta}_r$ is small and equations

(58) and (59) hold in most cases when the α_j 's are 1; see Bixel *et al.* (1963). That being so, it is plausible that equations (58) and (59) may hold for $\tilde{\eta}_r \neq 1$ provided that the α_j 's were equal. This, of course, would have to be explored.

2 Results

Our main concern in the following is the understanding of the characteristics of the derivative curve, $p'_{D_1}(\mathbf{x}_D; t_D)$, as these curves best demonstrate the characteristics of the rock under consideration. For each of the solutions presented, the pressure response, $p_D(\mathbf{x}_D; t_D)$, is also computed and examined. Our focus is responses at wells on the same side of the interface. Simulations shown were obtained through equation (36) through the aid of the Stehfest algorithm (1970a, b) as is normal for most well-test problems.

2.1 Perfect contact

Figure 1 presents derivative, $p'_{wD}(t_D)$, responses for a number of situations; the intent is to demonstrate the structure of the solutions and the influence of properties, specifically the exponent α_j and the ratio T_D . As one of our interests here is to demonstrate the accuracy of the solutions we assume the diffusivity ratio, η_D , to be 1. All properties considered are noted in Figure 1. The symbol L_D that is the dimensionless distance to the fault is defined by

$$L_D = \frac{L}{\ell}. \quad (60)$$

Simulations obtained through equation (36) for Υ_D of 0 are shown as unbroken lines. Because the properties of Zone 1 are held constant, all solutions originate from a single curve and branch out at later times as the result of the influence of Zone 2. Considering the specifics of the results shown here the bottommost curve labelled A and topmost curve labelled D correspond to values of T_D of 100 and 0.01, respectively, for $\alpha_j = 0.9$; $j = 1, 2$ with L_D of 100; the long-time approximations shown as dashed lines (marked AN) for identical conditions are in excellent agreement as they meet the stipulations in equation (59). Curves A and D may also be considered proxies for constant pressure and closed boundaries, respectively. Curves B and C correspond to α_2 of 0.7 and 0.5, respectively, each for T_D of 100; all three curves, A, B and C deviate from each other as the influence of Zone 2 becomes dominant. Not unexpectedly, long-time responses reflect a system of lower effective permeability as α_2 is smaller. It is interesting that Curves C and D essentially merge at long enough times. This result, as usual, suggests the difficulties one encounters in addressing the uniqueness of results obtained by evaluating the pressure-time curve; not unexpectedly supplementary information helpful (see Sect. 4). In Figure 2 we continue the exploration of the influence of properties, specifically the influences of η_D and α_j . The solutions are for T_D of 100 and α_1 of 0.9. The top set corresponds to dimensionless pressure, p_{wD} , response, and the bottom set represents the corresponding derivative, p'_{wD} . Again the properties of

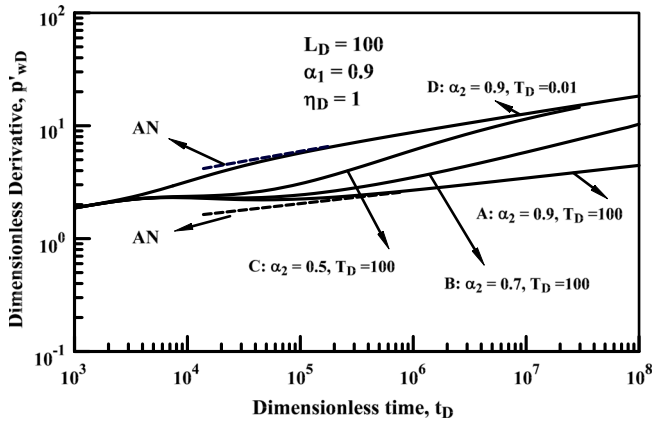


Fig. 1. Responses at a well producing a composite system under subdiffusive flow. All solutions generated through equation (36) for Y_D of 0 illustrate derivative responses, p'_{wD} . The influences of T_D , and the exponent, α_j , are considered; the properties of Zone 1 are held constant. The two dashed lines marked AN reflect calculations obtained through equation (59) and are in excellent agreement with the rigorous solutions which meet the two stipulations (equal $\tilde{\eta}_j$'s and α_j 's) stated. The upward rise in the responses as the magnitude of the exponent α_2 decreases reflects the lower effective permeability of Zone 2.

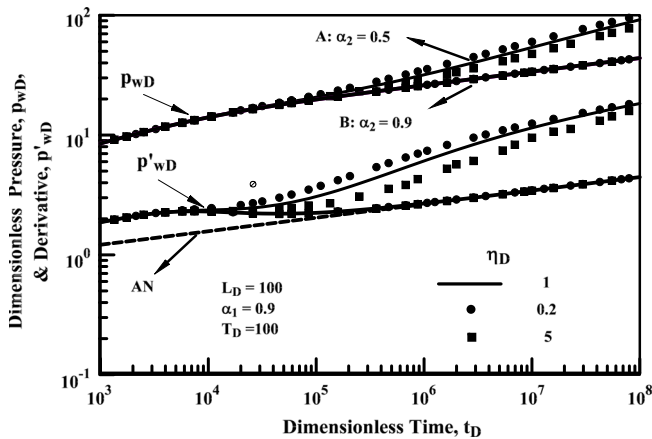


Fig. 2. The well response for production at a constant rate for a composite system; pressure, p_{wD} , and the corresponding derivative, p'_{wD} , responses, the top and bottom sets of curves, respectively, are shown. The intent is to demonstrate the influence of diffusivity ratio, η_D , and the exponent, α_2 . The unbroken lines presume that the η_D 's are equal; the symbols denoted as circles and squares presume that $\eta_D \neq 1$. If the α_j 's are equal, then the influence of η_D is negligibly small; see Curves labelled B; the corresponding derivative curve shows that the analytical solution shown as a dashed line (marked AN) predicts values in excellent agreement. The circles and squares illustrate the influence of η_D for $\alpha_2 = 0.5$. In general, solutions depend on both η_D and α_D .

Zone 1 are held constant; thus all solutions at later times emanate from a single curve, the branching depends on η_D and/or α_2 as soon as the region beyond the interface influences the well response. As in Figure 1, we assume

L_D to be 100. The dashed line (AN) reflects derivative responses given in equation (59) when both the diffusivities, η 's and α 's are equal. Two sets of solutions, Set A and Set B, are considered for α_2 of 0.9 and 0.5 and three values of η_D : 5 (squares), 1 (unbroken lines) and 0.2 (squares). Considering the $p_{wD}(t_D)$ solutions for α_2 of 0.5, namely Set A, we see that the influence of η_D is small – simulations for $\eta_D = 0.2$, the circles, and that for $\eta_D = 5$, the squares, straddle the $\eta_D = 1$ solution. If the α_j 's are equal (0.9 case), however, the influence of η_D is insignificant, Set B. These observations carry over to $p'_{wD}(t_D)$ curves. In essence, these results, and others not discussed here, show that the Bixel *et al.* (1963) observation that the influence of η_D is insignificant carries over to the subdiffusive solutions but the α_j 's must be equal.

2.2 Influence of a contact resistance

Figure 3 presents the influence of a fault, $\Upsilon_D = 10^3$, located at a distance of L_D of 100 at an observation point, $x_D = 60$, $y_D = 40$ with both the well and the observation location on the same side of the interface. The influences of T_D , η_D and α_2 for fixed values of Zone 1 properties with α_1 of 1, which result in all curves starting from the same point, are presented. All solutions except the line marked as Label A are for η_D of 8; the one identified as A is for an $\eta_D = 1$. Dimensionless times in Figure 3 are normalized with respect $L_D = 100$; that is, all solutions shown here are expressed as t_D/L_D^2 . The unbroken lines identical to the ones given in Raghavan and Ozkan (2011) correspond to the classical case with the variable of interest being T_D ; we consider three values of T_D : 10, 1 and 0.5. The peak in the derivative curve, not unexpectedly, reflects the restriction to flow as a result of the fault, with the slight concave upward behavior being the influence of the ratio $(\phi c_t)_1/(\phi c_t)_2$. Ultimately all lines become flat and reflect $2/(1 + T_D)$ as both α 's are 1 (see Eq. (59)).

The situation is different if $\alpha_2 < 1$; the dashed lines in Figure 3 reflect solutions for α_2 of 0.9 for identical values of T_D . These solutions deviate from the unbroken lines with the upward trend representing the fact of an additional restriction, namely that α_2 is now 0.9. A further reduction in α_2 will result in an earlier deviation and the magnitude of the upward trend would be larger; see line with small dashes for T_D of 10 which represents α_2 of 0.8. It is clear that subdiffusive characteristics of Zone 2 are reflected across the fault.

Earlier the influence of the porosity-compressibility ratio was addressed and it was noted that the concave upward portion was reflective of the value of η_D . The basis of this observation may be understood by comparing the curve labelled A which corresponds to $\eta_D = 1$ and $T_D = 1$ with its counterpart the solution for $\eta_D = 8$ and $T_D = 1$.

Results that are the complement of the ones in Figure 3 are discussed in Figure 4. Incidentally, here, we also draw attention to particular features of subdiffusion. The unbroken lines are responses for $\alpha_1 = 0.9$ and $\alpha_2 = 1.0$; that is, subdiffusive flow in Zone 1 and classical flow in Zone 2 – the opposite of situations considered in Figure 3 although a slightly different value of η_D is used and the range of T_D values is larger. While comparing the responses

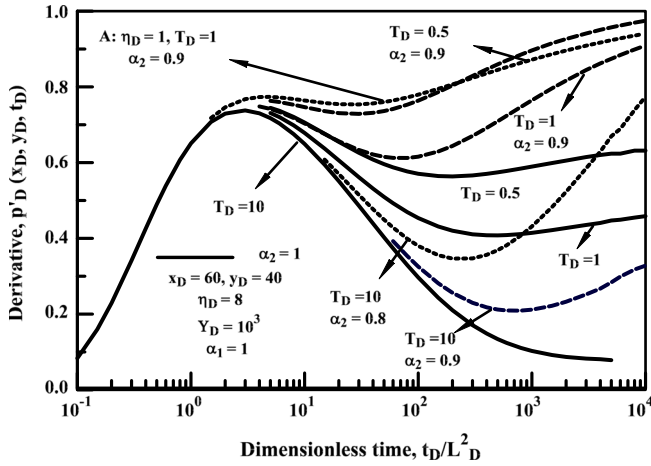


Fig. 3. Influence of contact resistance on derivative responses at an observation well, $Y_D = 10^3$. Both wells are on the same side of the fault, hence the sharp rise in the derivative curve. The principal goal is to demonstrate the influence of T_D , and α_2 . Properties of Zone 1 are held constant and classical diffusion prevails ($\alpha_1 = 1.0$); unbroken lines are for $\alpha_2 = 1$; dashed lines for $\alpha_2 = 0.9$ and small dashes for $\alpha_2 = 0.8$.

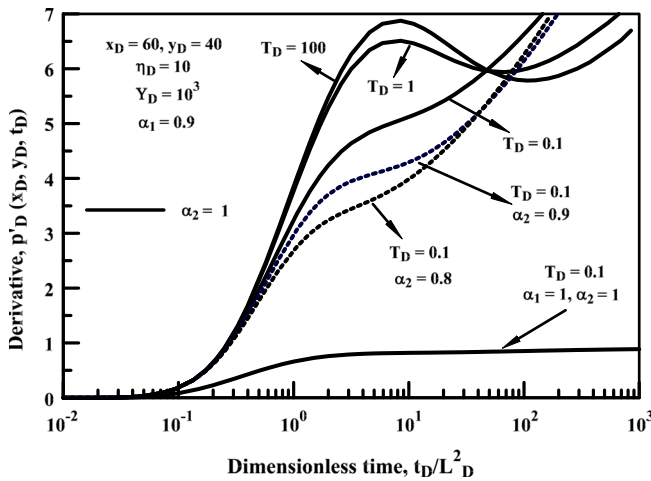


Fig. 4. Influence of contact resistance on derivative responses at an observation well, $Y_D = 10^3$. Again, both wells are on the same side of the fault, the change in α_1 from 1.0 to 0.9 produces a significant change in the magnitude of the response.

in Figures 3 and 4 the significant change in the values of the ordinate needs to be borne in mind because a change in the magnitude of α_1 from 1.0 is 0.9 results in a significant change in the magnitude of $p_{wD}(t_D)$. All solutions considered here except for the bottommost curve assume subdiffusive flow in Zone 1, and, again all solutions for α_2 of 0.9 start out together with the steep rise reflecting the conductivity of the fault; it is clear that subdiffusive effects of Zone 1 are evident even when the properties of Zone 2 are dominant unlike what we would expect if α_1 were 1.0 as shown by the bottommost curve in Figure 3. The two dashed lines presume subdiffusive flow in Zone 2; initially responses fall

below that for α_2 of 1.0 reflecting the speed of propagation as discussed in Section 1. With time, however, the slopes of the curves become steeper reflecting a lower effective permeability as α_2 decreases. In practice such steeper slopes would not reflect additional boundary effects.

3 Discussion

Before closing with a few general observations, for practicality, we note that the exponent α may be rendered as a Probability Density Function (PDF) through the instantaneous Green's function, $G(r, t)$, through the expression in equation (18). The first term yields:

$$\mathcal{G}(\mathbf{r}; t) = \frac{1}{4\pi\tilde{\eta}t^\alpha} H_{1,2}^{2,0} \left[\frac{r^2}{4\tilde{\eta}t^\alpha} \middle|_{(0,1),(0,1)}^{1-\alpha,\alpha} \right], \quad (61)$$

where r is the distance from the source. By considering a series of transformations of the Fox function Schneider and Wyss (1989) arrive at an expression for $\mathcal{G}(\mathbf{r}; t)$ in the following form,

$$\mathcal{G}(\mathbf{r}; t) \sim \frac{\alpha^{(d\alpha-1)/2}}{\sqrt{2-\alpha}} (4\pi\tilde{\eta}t^\alpha)^{-d/2} \left(\frac{r^2\alpha^\alpha}{4\tilde{\eta}t^\alpha} \right)^{-\frac{d(1-\alpha)}{2(2-\alpha)}} \times \exp \left[-(2-\alpha) \left(\frac{r^2\alpha^\alpha}{4\tilde{\eta}t^\alpha} \right)^{\frac{1}{(2-\alpha)}} \right], \quad (62)$$

where the symbol d is the number of dimensions; that is, equation (62) addresses a far more general problem than the one we consider for $d = 2$ in our case. $\mathcal{G}(\mathbf{r}; t)$ exhibits power-law behavior at long-enough times for $\alpha < 1$. This expression agrees with those in Carslaw and Jaeger (1959) for $d \leq 3$ if $\alpha = 1$. Grebenkov (2010) arrives at the same expression as that given in equation (62) but for $\mathcal{G}(\mathbf{r}; t \rightarrow 0)$ by using the result

$$L^{-1}[s^{-\kappa} \exp(-r\sqrt{s^2/\tilde{\eta}})] \simeq \frac{\alpha^{\kappa+1/2}}{\sqrt{\pi(2-\alpha)}} \left(\frac{r^2\alpha^\alpha}{4\tilde{\eta}t^\alpha} \right)^{\frac{(\kappa+1/2)}{(2-\alpha)}} t^{-(\kappa+1)} \times \exp \left[-(2-\alpha) \left(\frac{r^2\alpha^\alpha}{4\tilde{\eta}t^\alpha} \right)^{\frac{1}{(2-\alpha)}} \right], \quad (63)$$

with $\kappa = (d + 1)/4 - 1$. Incidentally, the left-hand side of equation (63) may also be expressed in terms of the Wright function, $W(\alpha, \beta; t)$, as indicated in Povstenko (2015) by

$$L^{-1}[s^{-\beta} \exp(-\lambda s^\alpha)] = t^{\beta-1} W(-\alpha, \beta; -\lambda t^{-\alpha}), \quad (64)$$

$$0 < \alpha < 1, \lambda > 0.$$

The moments of an even order which would also be of interest are (Uchaikin, 2013); see equation (1):

$$\int_{\mathcal{R}^d} |\mathbf{x}|^{2n} \mathcal{G}(|\mathbf{x}|, t) d\mathbf{x} = \frac{\Gamma(n + d/2)\Gamma(n + 1)}{\Gamma(d/2)\Gamma(\alpha n + 1)} \eta t^{n\alpha}, \quad (65)$$

$$n = 0, 1, 2, \dots$$

Variations of $\mathcal{G}(\mathbf{r}; t)$ as a function of r are shown in Figure 5 for three values of α : 1, 0.9, and 0.8. These values were

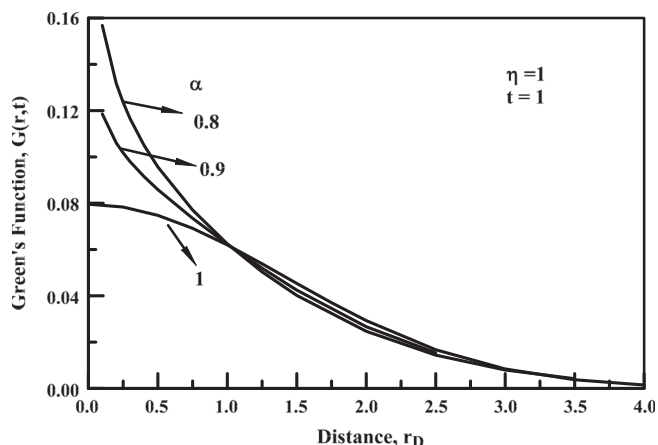


Fig. 5. The probability density function as a function of α in Figure 5; $\eta = 1$ and $t = 1$.

computed independent of the expressions (62) and (63). The characteristic bell shape is evident only for $\alpha = 1$. Note that $\mathcal{G}(r; t)$ may be estimated by equation (62) for $r = 0$ for α of 1.

4 Commentary

Although our goal is to place a focus on the mathematical aspects and characteristics of the responses under subdiffusive flow, we briefly mention a few practical aspects we follow. Evaluation of well responses is to be done in the same manner as that for classical diffusion. The additional complication is the determination of the exponents α_j (or at least α_1) which would be the first step in identifying subdiffusive flow; the determination of the exponent would depend, among other things, on the duration of the test; for in the case of other factors such as the properties of the rock or well locations we may not have much control. The exponent is best determined through the derivative curve, $\Delta p'$, as this curve when combined with the pressure response, Δp , better constrains the “uncertainty in the solution space”; see Raghavan (2004). It is for this reason we have, as mentioned, discussed the derivative extensively in the text. Given that the process is to be data driven we should garner as much information as possible regarding reservoir architecture (maps, cores, logs, etc.); see Raghavan *et al.* (2001) for a comprehensive review of combining, geology, geophysics and production statistics. In addition, it is essential to know completion information; see Raghavan (1993). We should note that although the analysis is conducted on the Theis (1935) scale, the use of a subdiffusive model does not imply a single value for properties. The end product is to arrive at a reservoir description that reproduces future performance. For all of this the derivative curve should be the starting point.

5 Concluding remarks

A 2D source function solution under subdiffusion in two contiguous media separated by a partially absorbing or

communicating interface is documented. In all other aspects the media are infinite in extent. Two dimensional subdiffusion problems, to our knowledge, are yet to be considered. The role of the interface exerting a resistance is detailed. For the situation described exact and asymptotic solutions are derived. Computational results are discussed in the context of applications to the earth sciences; specifically the solutions may be applied to hydrocarbon or geothermal-fluid extraction, flow of groundwater, roles of faults, and sequestration. Extensions to other situations related to flow in porous media are readily possible: variable-rate problems through Duhamel’s theorem including situations involving wellbore storage effects, multiple, contiguous media, complex wellbores such as inclined, fractured and horizontal wells and equivalent porous media models such as those described in Barenblatt *et al.* (1960) and de Swaan-O (1976). Quantitative descriptions of subdiffusive flow and their effects on estimating properties of the rock and permeable interfaces are discussed. Should subdiffusive effects exist it is not possible to ignore the existence of their time-dependent influences at long enough times as in the case of classical diffusion.

In summary, the Green’s function formulation to address the role of a partially communicating interface detailed here is a general one; we may tap into the potential of the solutions discussed here in a number of contexts such as biophysics, fractal media, and the like, as Laplace transformations and the method of sources and sinks are used. As mentioned in the Introduction, the expressions derived here are directly transferable to other fields; only a change in Nomenclature is needed.

References

- Albinali A., Holy R., Sarak H., Ozkan E. (2016) Modeling of 1D anomalous diffusion in fractured nanoporous media, *Oil Gas Sci. Technol.- Rev. IFP Energies nouvelles* **71**, 56.
- Angulo J.M., Ruiz-Medina M.D., Anh V.V., Grecksch W. (2000) Fractional diffusion and fractional heat equation, *Adv. Appl. Probab.* **32**, 4, 1077–1099.
- Barenblatt G.I., Zheltov YuP, Kochina I.N. (1960) Basic concepts in the theory of seepage of homogeneous liquids in fissured rocks, *J. Appl. Math. Mech.* **24**, 5, 1286–1303.
- Bense V., Person M. (2006) Faults as conduit-barrier system to fluid flow in siliclastic sedimentary aquifers, *Water Resour. Res.* **42**, W05421. doi: [10.1029/2005WR004480](https://doi.org/10.1029/2005WR004480).
- Bernard S., Delay F., Porel G. (2006) A new method of data inversion for the identification of fractal characteristics and homogenization scale from hydraulic pumping tests in fractured aquifers, *J. Hydrol.* **328**, 3, 647–658. doi: [10.1016/j.jhydrol.2006.01.008](https://doi.org/10.1016/j.jhydrol.2006.01.008).
- Bixel H.C., Larkin B.K., Van Poollen H.K. (1963) Effect of linear discontinuities on pressure build-up and drawdown behavior, *J. Pet. Tech.* **15**, 8, 885–895. doi: [10.2118/611-PA](https://doi.org/10.2118/611-PA).
- Caine J.S., Evans J.P., Forster C.B. (1996) Fault zone architecture and permeability structure, *Geology* **24**, 11, 1025–1028.
- Caputo M. (1967) Linear models of dissipation whose Q is almost Frequency Independent-II, *Geophys. J. R. Astron. Soc.* **13**, 5, 529–539.

- Caputo M. (1999) Diffusion of fluids in porous media with memory, *Geothermics* **28**, 1, 113–130.
- Carslaw H.S., Jaeger J.C. (1959) *Conduction of heat in solids*, 2nd edn., Vol. **22**, Clarendon Press, Oxford, pp. 319–326, 327–352.
- Chang J., Yortsos Y.C. (1990) Pressure-transient analysis of fractal reservoirs, *SPE Form. Eval.* **5**, 1, 31–39.
- Childs E.C. (1969) *An Introduction to the Physical Basis of Soil Water Phenomena*, John Wiley and Sons Ltd, London, pp. 153–178.
- Cinco-Ley H., Samaniego-V. F., Kuchuk F. (1985) The pressure transient behavior for naturally fractured reservoirs with multiple block size. In: *Paper SPE 14168 Presented at SPE Annual Technical Conference and Exhibition*, 22–26 September, Las Vegas, Nevada.
- Cooper H.H., Jacob C.E. (1946) A generalized graphical method for evaluating formation constants and summarizing well-field history, *Trans. AGU* **27**, 526–534.
- Cortis A., Knudby C. (2006) A continuous time random walk approach to transient flow in heterogeneous porous media, LBNL-59885, *Water Resour. Res.* **42**, W10201.
- Darcy H. (1856) *Les Fontaines publiques de la ville de Dijon, Exposition et application des principes à suivre et des formules à employer dans les questions de distribution d'eau ; ouvrage terminé par un appendice relatif aux fournitures d'eau de plusieurs villes au filtrage des eaux et à la fabrication des tuyaux de fonte, de plomb*, Victor Dalmont, Paris, pp. 590–594.
- Dassas Y., Duby Y. (1995) Diffusion toward fractal interfaces, potentiostatic, galvanostatic, and linear sweep voltammetric techniques, *J. Electrochem. Soc.* **142**, 12, 4175–4180.
- de Swaan-O A. (1976) Analytical solutions for determining naturally fractured reservoir properties by Well testing, *Soc. Pet. Eng. J.* **16**, 3, 117–122. doi: [10.2118/5346-PA](https://doi.org/10.2118/5346-PA).
- Erdélyi A., Magnus W., Oberhettinger F., Tricomi F.G. (1954) *Tables of integral transforms, based, in part, on notes left by Harry Bateman and Compiled by the Staff of the Bateman Manuscript Project*, vol. **1**, McGraw-Hill, New York, 17.
- Erdélyi A., Magnus W.F., Oberhettinger F., Tricomi F.G. (1955) *Higher transcendental functions*, vol. **3**, McGraw-Hill, New York, pp. 206–227.
- Evans J.P. (1988) Deformation mechanisms in granitic rocks at shallow crustal levels, *J. Struct. Geol.* **10**, 5, 437–443.
- Feller W. (1971) *An introduction to probability theory and its applications. II*, 2nd edn., Wiley, New York, pp. 8–10, 50.
- Fomin S., Chugunov V., Hashida T. (2011) Mathematical modeling of anomalous diffusion in porous media, *Fractional Differ. Calc.* **1**, 1–28.
- Fu L., Milliken K.L., Sharp J.M. Jr. (1994) Porosity and permeability variations in fractured and liesegang-banded Breathitt sandstones (Middle Pennsylvanian), eastern Kentucky: Diagenetic controls and implications for modeling dual-porosity systems, *J. Hydrol.* **154**, 1–4, 351–381.
- Gefen Y., Aharony A., Alexander S. (1983) Anomalous diffusion on percolating clusters, *Phys. Rev. Lett.* **50**, 1, 77–80. doi: [10.1103/PhysRevLett.50.77](https://doi.org/10.1103/PhysRevLett.50.77).
- Gradshteyn I.S., Ryzhik I.M. (1994) *Table of integrals, series and products*, 5th edn., Vol. **532**, Academic Press Inc., Orlando, pp. 3.961–3.962.
- Grebekov D.B. (2010) Subdiffusion in a bounded domain with a partially absorbing-reflecting boundary, *Phys. Rev. E: Stat. Phys. Plasmas Fluids* **81**, 1, 021128.
- Gurtin M.E., Pipkin A.C. (1968) A general theory of heat conduction with finite wave speeds, *Arch. Rational Mech. Anal.* **31**, 2, 113–126. doi: [10.1007/BF00281373](https://doi.org/10.1007/BF00281373).
- Hänninen J.J., Pirjola R.J., Lindell I.V. (2002) Application of the exact image theory to studies of ground effects of space weather, *Geophys. J. Int.* **151**, 2, 534–542.
- Henry B.I., Langlands T.A.M., Straka P. (2010) An introduction to fractional diffusion, in: Presented at the Conference: Complex Physical, Biophysical and Econophysical Systems – Proceedings of the 22nd Canberra International Physics Summer School, pp. 37–89.
- Hilfer R., Anton L. (1995) Fractional master equations and fractal time random walks, *Phys. Rev. E* **51**, 2, R848–R851.
- Jourde H., Pistrea S., Perrochet P., Droguea C. (2002) Origin of fractional flow dimension to a partially penetrating well in stratified fractured reservoirs, new results based on the study of synthetic fracture networks, *Adv. Water Res.* **25**, 4, 371–387.
- Kenkre V.M., Montroll E.W., Shlesinger M.F. (1973) Generalized master equations for continuous-time random walks, *J. Stat Phys* **9**, 1, 45–50.
- Kim S., Kavvas M.L., Ercan A. (2015) Fractional ensemble average governing equations of transport by time-space nonstationary stochastic fractional advective velocity and fractional dispersion, II: Numerical investigation, *J. Hydrol. Eng.* **20**, 2, 04014040.
- Koztoowicz K. (2017) Subdiffusion in a system consisting of two different media separated by a thin membrane, *Int. J. Heat Mass Transfer* **111**, 1322–1333.
- Lê Mehautê A., Crepy G. (1983) Introduction to transfer and motion in fractal media: The geometry of kinetics, *Solid State Ion.* **1**, 9–10, 17–30.
- Lindell I., Alanen E. (1984) Exact image theory for the Sommerfeld half-space problem, part III: General formulation, *IEEE Trans. Antennas Propag.* **32**, 10, 1027–1032. doi: [10.1109/TAP.1984.1143204](https://doi.org/10.1109/TAP.1984.1143204).
- Lindell I., Alanen E., von Bagh H. (1986) Exact image theory for the calculation of fields transmitted through a planar interface of two media, *IEEE Trans. Antennas Propag.* **34**, 2, 29–137. doi: [10.1109/TAP.1986.1143788](https://doi.org/10.1109/TAP.1986.1143788).
- Lindquist E. (1933) On the flow of water through porous soil, *1st Congres des Grands Barrages, Stockholm*, **5**, 91–101.
- Magin R.L., Ingo C., Colon-Perez L., Triplett W., Mareci T.H. (2013) Characterization of anomalous diffusion in porous biological tissues using fractional order derivatives and entropy, *Microporous Mesoporous Mater.* **178**, 15, 39–43. doi: [10.1016/j.micromeso.2013.02.054](https://doi.org/10.1016/j.micromeso.2013.02.054).
- Mainardi F. (2010) *Fractional calculus and waves in linear viscoelasticity*, Imperial College Press, London, pp. 211–236.
- Mainardi M., Pagnini G., Saxena R.K. (2005) Fox H functions in fractional diffusion, *J. Comput. Appl. Math.* **178**, 1–2, 321–331.
- Mandelis A., Nicolaidis L., Chen Y. (2001) Structure and the reflectionless/refractionless nature of parabolic diffusion-wave fields, *Phys. Rev. Lett.* **87**, 2, 020801. doi: [10.1103/PhysRevLett.87.020801](https://doi.org/10.1103/PhysRevLett.87.020801).
- Mathai A.M., Saxena R.K. (1978) *The H-function with applications in statistics and other disciplines*, Wiley, New Delhi, India, pp. 1–19.
- Metzler R., Glockle W.G., Nonnenmacher T.F. (1994) Fractional model equation for anomalous diffusion, *Physica A* **211**, 1, 13–24.

- Miller B. (2005) The Baton Rouge Fault: Conduit or impediment to groundwater flow? in: *Paper Presented at 54th Annual Meeting, Southeast*, Sect. of the Geol. Soc. of Am., Biloxi Miss.
- Miller F.G. (1954) Multiphase flow theory and the problem of spacing oil wells, *United States Bureau of Mines* **529**, 8–10.
- Mitchell T.M., Faulkner D.R. (2009) The nature and origin of off-fault damage surrounding strike-slip fault zones with a wide range of displacements: A field study from the Atacama fault system, northern Chile, *J. Struct. Geol.* **31**, 8, 802–816. doi: [10.1016/j.jsg.2009.05.002](https://doi.org/10.1016/j.jsg.2009.05.002).
- Molz F.J. III, Fix G.J. III, Lu S.S. (2002) A physical interpretation for the fractional derivative in Levy diffusion, *Appl. Math. Lett.* **15**, 7, 907–911.
- Montroll E.W., Weiss G.H. (1965) Random walks on lattices II, *J. Math. Phys.* **6**, 167–181.
- Moodie T.B., Tait R. (1983) On thermal transients with finite wave speeds, *J. Acta Mech.* **50**, 1–2, 97–104. doi: [10.1007/BF01170443](https://doi.org/10.1007/BF01170443).
- Neville I.R., Sharp, Jr. J.M., Kreisel I. (1998) Contaminant transport in sets of parallel finite fractures with fracture skins, *J. Contam. Hydrol.* **31**, 1–2, 83–109.
- Nigmatullin R. (1984a) On the theory of relaxation with “remnant” memory, *Phys. Stat. Sol. B* **124**, 1, 389–393. doi: [10.1002/pssb.2221240142](https://doi.org/10.1002/pssb.2221240142).
- Nigmatullin R.R. (1984b) To the theoretical explanation of the universal response, *Phys. Status Solidi B Basic Res.* **123**, 2, 739–745.
- Noetinger B., Estebenet T. (2000) Up-scaling of double porosity fractured media using continuous-time random walks methods, *Transp. Porous Med.* **39**, 3, 315–337.
- Norwood F.R. (1972) Transient thermal waves in the general theory of heat conduction with finite wave speeds, *ASME J. Appl. Mech.*, **39**, 3, 673–676. doi: [10.1115/1.3422771](https://doi.org/10.1115/1.3422771).
- Oberhettinger F., Badii L. (1973) *Tables of Laplace transforms*, Springer Verlag, Berlin, p. 268.
- Odling N.E., Gillespie P., Bourguine B., Castaing C., Chiles J.P., Christiansen N.P., Fillion E., Albert G., Olsen C., Lena T., Trice R., Aarseth E.S., Walsh J.J., Watterson J. (1999) Variations in fracture system geometry and their implications for fluid flow in fractured hydrocarbon reservoirs, *Petrol. Geosci.* **5**, 373–384.
- O’Shaughnessy B., Procaccia I. (1985a) Analytical solutions for diffusion on fractal objects, *Phys. Rev. Lett.* **54**, 5, 455–458.
- O’Shaughnessy B., Procaccia I. (1985b) Diffusion on fractals, *Phys. Rev. A: At. Mol. Opt. Phys* **32**, 5, 3073–3083.
- Philip J.R. (1957) Transient fluid motions in saturated porous media, *Aust. J. Phys.* **10**, 43–53.
- Plöciniczak Ł. (2015) Analytical studies of a time-fractional porous medium equation. Derivation, approximation and applications, *Commun. Nonlinear Sci. Numer. Simul.* **24**, 1–3, 169–183.
- Povstenko Y. (2015) Linear fractional diffusion-wave equation for scientists and engineers, *Birkhäuser* **24–30**, 32, 34.
- Povstenko Y.Z. (2013) Fractional heat conduction in infinite one-dimensional composite medium, *J. Therm. Stresses* **36**, 4, 351–363. doi: [10.1080/01495739.2013.770693](https://doi.org/10.1080/01495739.2013.770693).
- Povstenko Y., Kyrylych T. (2019) Fractional heat conduction in solids connected by thin intermediate layer: Nonperfect thermal contact, *Continuum Mech. Thermodyn.* **31**, 1719–1731, doi: [10.1007/s00161-019-00750-w](https://doi.org/10.1007/s00161-019-00750-w).
- Prats M., Raghavan R. (2013) Finite horizontal well in a uniform-thickness reservoir crossing a natural fracture, *SPE J.* **18**, 5, 982–992. doi: [10.2118/163098-PA](https://doi.org/10.2118/163098-PA).
- Prats M., Raghavan R. (2014) Properties of a natural fracture and its skins from reservoir well tests, *SPE J.* **19**, 3, 390–397. doi: [10.2118/167262-PA](https://doi.org/10.2118/167262-PA).
- Raghavan R. (1993) *Well test analysis*, Prentice Hall, Englewoods Cliffs, NJ, pp. 6–8, 13.
- Raghavan R. (2004) A review of applications to constrain pump-test responses to improve on geological description and uncertainty, *Rev. Geophys.* **42**, 1–29.
- Raghavan R. (2010) A composite system with a planar interface, *J. Pet. Sci. Eng.* **70**, 3–4, 229–234. doi: [10.1016/j.petrol.2009.11.015](https://doi.org/10.1016/j.petrol.2009.11.015).
- Raghavan R. (2011) Fractional derivatives: Application to transient flow, *J. Pet. Sci. Eng.* **80**, 1, 7–13. doi: [10.1016/j.petrol.2011.10.003](https://doi.org/10.1016/j.petrol.2011.10.003).
- Raghavan R. (2012) A horizontal well in a composite system with planar interfaces, *Adv. Water Res.* **38**, 38–47. doi: [10.1016/j.advwatres.2011.12.009](https://doi.org/10.1016/j.advwatres.2011.12.009).
- Raghavan R., Chen C. (2017) Addressing the influence of a heterogeneous matrix on well performance in fractured rocks, *Transp. Porous Med.* **117**, 1, 69–102. doi: [10.1007/s11242-017-0820-5](https://doi.org/10.1007/s11242-017-0820-5).
- Raghavan R., Ozkan E. (1994) *A method for computing unsteady flows in porous media*, Pitman Research Notes in Mathematics Series (318), Longman Scientific & Technical, Harlow, UK, p. 188.
- Raghavan R., Ozkan E. (2011) Flow in composite slabs. *SPE Journal* **16**, 2, 374–387. doi: [10.2118/140748-PA](https://doi.org/10.2118/140748-PA).
- Raghavan R., Chen C. (2018) Time and space fractional diffusion in finite systems, *Transp. Porous Med.* **123**, 173–193.
- Raghavan R., Chen C. (2019) The Theis solution for subdiffusive flow in rocks, *Oil Gas Sci. Technol. - Rev. IFP Energies nouvelles* **74**, 6. doi: [10.2516/ogst/2018081](https://doi.org/10.2516/ogst/2018081).
- Raghavan R., Dixon T.N., Phan V.Q., Robinson S.W. (2001) Integration of geology, geophysics, and numerical simulation in the interpretation of a well test in a fluvial reservoir, *SPE Reserv. Evalu. Eng.* **4**, 3, 201–208. doi: [10.2118/72097-PA](https://doi.org/10.2118/72097-PA).
- Savage H.M., Brodsky E.E. (2011) Collateral damage: Evolution with displacement of fracture distribution and secondary fault strands in fault damage zones, *J. Geophys. Res.* **116**, B03405. doi: [10.1029/2010JB007665](https://doi.org/10.1029/2010JB007665).
- Saxena R.K., Mathai A.M., Haubold H.J. (2006) Fractional reaction-diffusion equations, *Astrophys. Space Sci.* **305**, 3, 289–296.
- Schneider W.R., Wyss W. (1989) Fractional diffusion and wave equations, *J. Math. Phys.* **30**, 134. doi: [10.1063/1.528578](https://doi.org/10.1063/1.528578).
- Scholz C.H., Dawers N.H., Yu J.Z., Anders M.H., Cowie P.A. (1993) Fault growth and fault scaling laws: Preliminary results, *J. Geophys. Res.* **98**, B12, 21951–21961.
- Sharp J.M. Jr., Kreisel I., Milliken K.L., Mace R.E., Robinson N.I. (1996) Fracture skin properties and effects on solute transport: Geotechnical and environmental implications, in: M. Aubertin, F. Hassam, H. Mitri (eds), *Rock Mechanics, Tools and Techniques*, Balkema, Rotterdam, pp. 1329–1335.
- Shendeleva M.L. (2004) Instantaneous line heat source near a plane interface, *J. Appl. Phys.* **95**, 5, 2839–2845.
- Sommerfeld A. (1909) Über die Ausbreitung der Wellen in der drahtlosen Telegraphie, *Ann. Phys.* **28**, 665–736.

- Stehfest H. (1970a) Algorithm 368: Numerical inversion of Laplace transforms [D5], *Commun. ACM* **13**, 1, 47–49.
- Stehfest H. (1970b) Remark on algorithm 368: Numerical inversion of Laplace transforms, *Commun. ACM* **13**, 10, 624.
- Su N., Nelson P.N., Connor S. (2015) The distributed-order fractional diffusion-wave equation of groundwater flow: Theory and application to pumping and slug tests, *J. Hydrol.*, **529**, 1262–1273. doi: [10.1016/j.jhydrol.2015.09.033](https://doi.org/10.1016/j.jhydrol.2015.09.033).
- Su N. (2014) Mass-time and space-time fractional partial differential equations of water movement in soils: Theoretical framework and application to infiltration, *J. Hydrol.* **519** (B), 1792–1803. doi: [10.1016/j.jhydrol.2014.09.021](https://doi.org/10.1016/j.jhydrol.2014.09.021)
- Suzuki A.T., Hashida K.L., Horne R.N. (2016) Experimental tests of truncated diffusion in fault damage zones, *Water Resour. Res.* **52**, 8578–8589. doi: [10.1002/2016WR019017](https://doi.org/10.1002/2016WR019017).
- Tao S., Gao X., Li C., Zeng J., Zhang X., Yang C., Zhang J., Gong Y. (2016) The experimental modeling of gas percolation mechanisms in a coal-measure tight sandstone reservoir: A case study on the coal-measure tight sandstone gas in the Upper Triassic Xu-Jiahe Formation, Sichuan Basin, China, *J. Nat. Gas Geosci.* **1**, 6, 445–455. doi: [10.1016/j.jnggs.2016.11.009](https://doi.org/10.1016/j.jnggs.2016.11.009).
- Theis C.V. (1935) The relationship between the lowering of the piezometric surface and the rate and duration of discharge of a well using ground-water storage, *Eos Trans. AGU* **2**, 519–524.
- Tualle J.-M., Tinet E., Prat J., Avriillier S. (2000) Light propagation near turbid-turbid planar interfaces, *Opt. Commun.* **183**, 5–6, 337–346.
- Uchaikin V.V. (2013) *Fractional derivatives for physicists and engineers. Volume I: Background and Theory*, Springer, New York, 151, 296.
- Yanga S., Zhoua H.W., Zhang S.Q., Ren W.G. (2019) A fractional derivative perspective on transient pulse test for determining the permeability of rocks, *Int. J. Rock Mech. Min. Sci.* **113**, 92–98.
- Yaxley L.M. (1987) Effect of a partially communicating fault on transient pressure behavior, *SPE Form. Eval.* **2**, 4, 590–598. doi: [10.2118/14311-PA](https://doi.org/10.2118/14311-PA).
- Zhokh A., Strizhak P. (2018) Non-Fickian transport in porous media: Always temporally anomalous? *Transp. Porous Med.* **124**, 2, 309–323. doi: [10.1007/s11242-018-1066-6](https://doi.org/10.1007/s11242-018-1066-6).
- Zhokh A., Strizhak P. (2019) Investigation of the anomalous diffusion in the porous media: A spatiotemporal scaling, *Heat Mass Transfer.* **55**, 1–10. doi: [10.1007/s00231-019-02602-4](https://doi.org/10.1007/s00231-019-02602-4).